

## Introduction

Geophysical imaging: to do what?

Generalities on Inverse Problems

## Seismic data

A first glance at seismic inversion methods

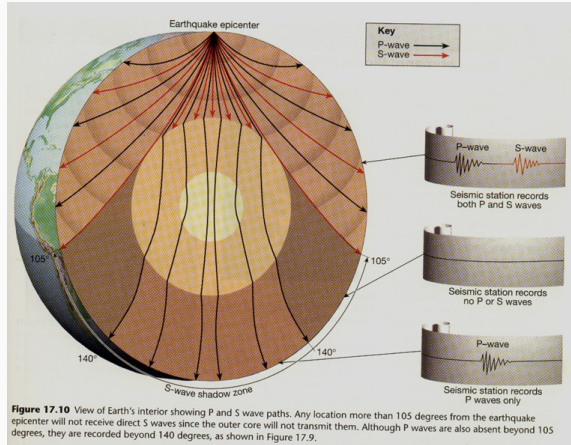
## Full waveform modeling

Building the wave equation

Heterogeneity, anisotropy and attenuation

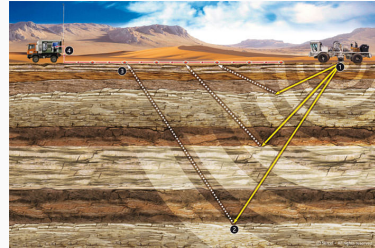
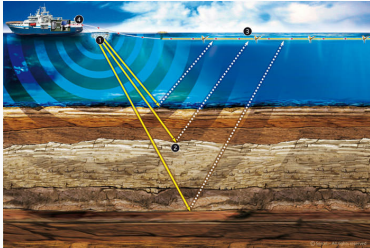
## Full waveform inversion

# We can use Earthquakes



*Global tomography sketch: an earthquake acts as a source which propagates elastic waves which are recorded by seismic stations spread at different point of the surface*





*Controlled source acquisition sketch, in a marine environment (left) or on land (right)*

In terms of mathematics, the seismic data is thus a collection of time functions  $d(t)$  associated with a source  $s$  and a receiver  $r$ . We will denote it as

$$d_{r,s}(t), \tag{2}$$

in the following, or equivalently

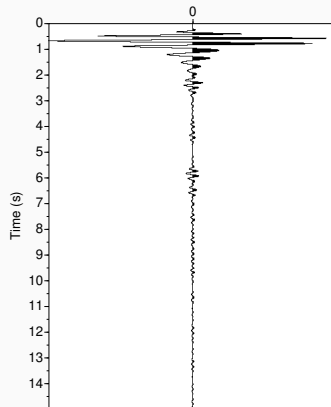
$$d(x_s, x_r, t), \tag{3}$$

or

$$d_s(x_r, t), \tag{4}$$

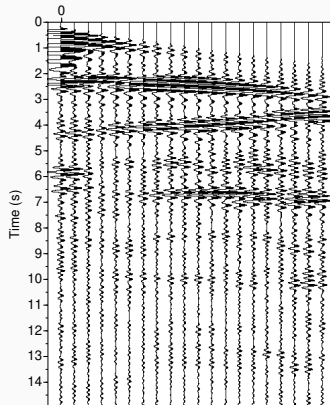
depending on the context. In these notations  $x_r$  and  $x_s$  denote the spatial position of the receiver  $r$  and the source  $s$  respectively. A single function  $d_{r,s}(t)$  will be referred to as a *seismic trace* in the following.

A typical example of a seismic trace is presented in Figure 20.



Typical seismic trace  $d(t)$  as a function of time.

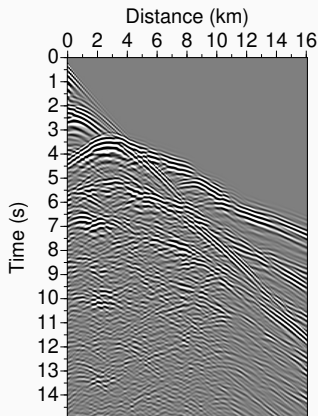
Instead of analyzing the data trace by trace: look simultaneously at several traces.



20 seismic traces  $d_r(t)$  as a function of time, depending on the receiver/source distance, also referred to as *offset* in the following.

# Seismogram

When the number of traces is even larger: use a 2D plot with a black & white chart

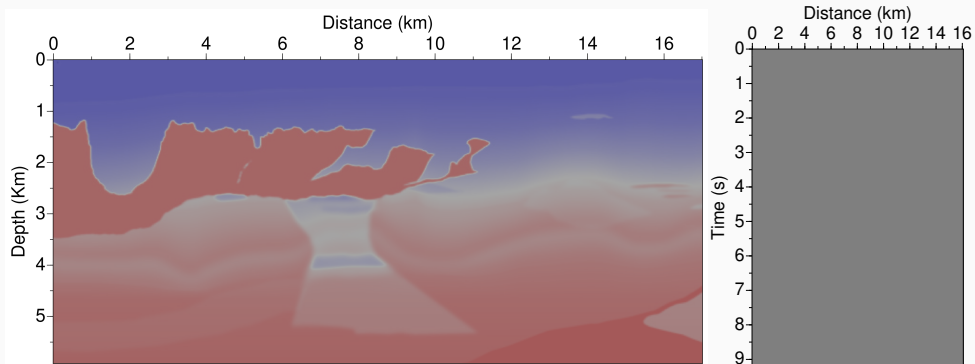


A typical seismogram in black and white representation. 161 traces spanning 16 km are used here.

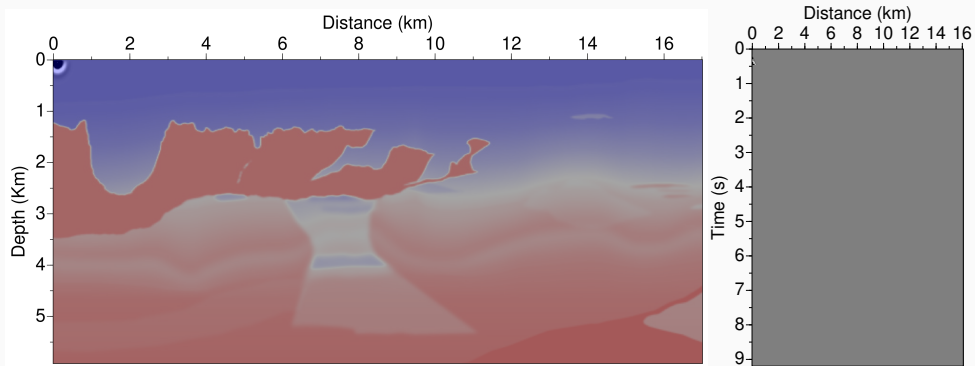
White correspond to negative values, black to positive values, while gray corresponds to 0. This yields

the typical seismogram representation, widely used in exploration geophysics.

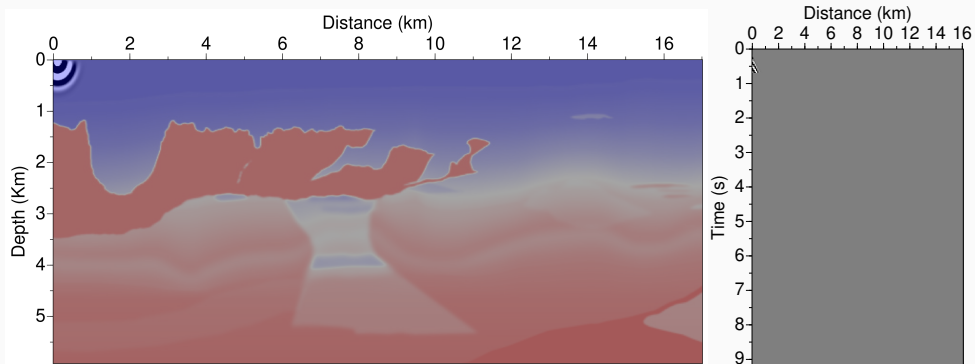
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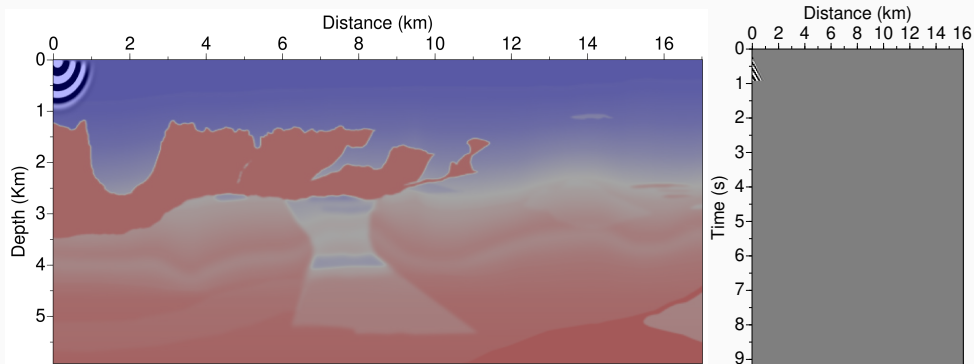


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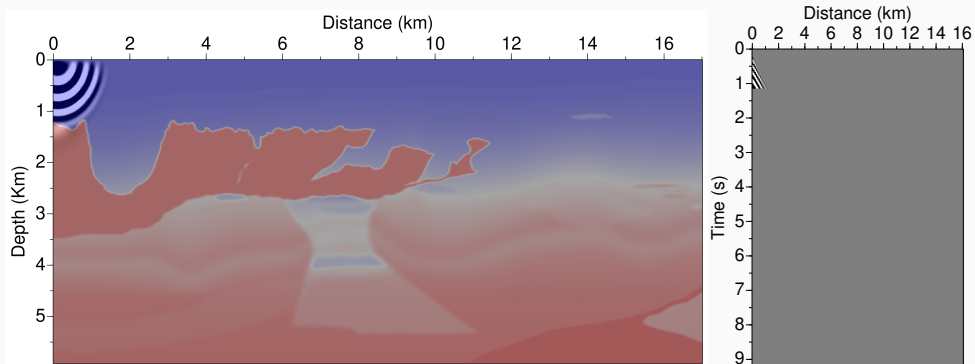




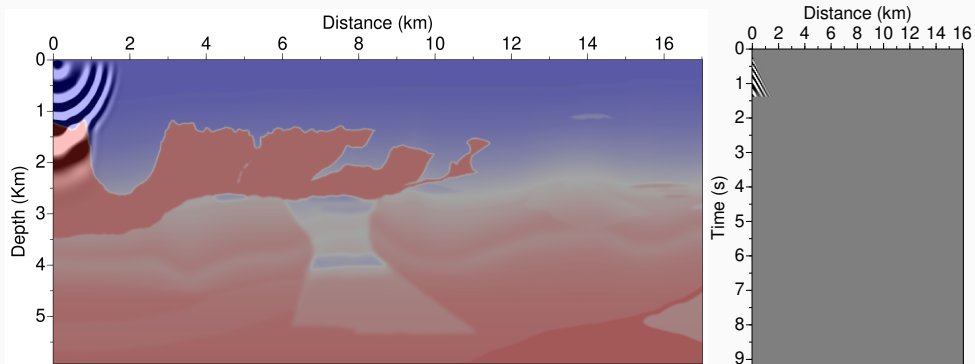
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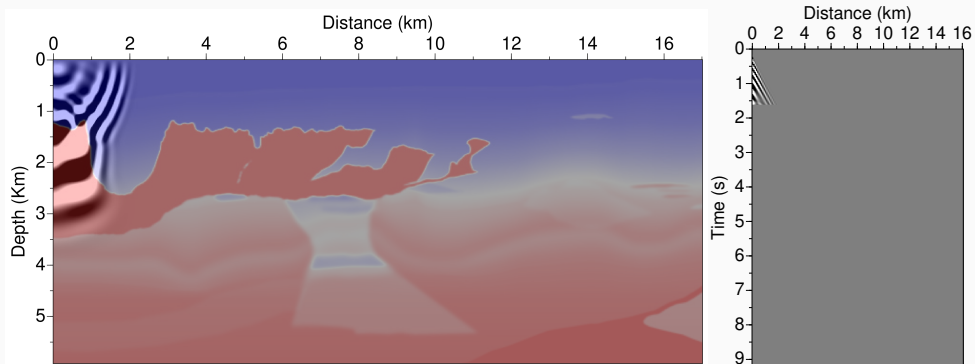
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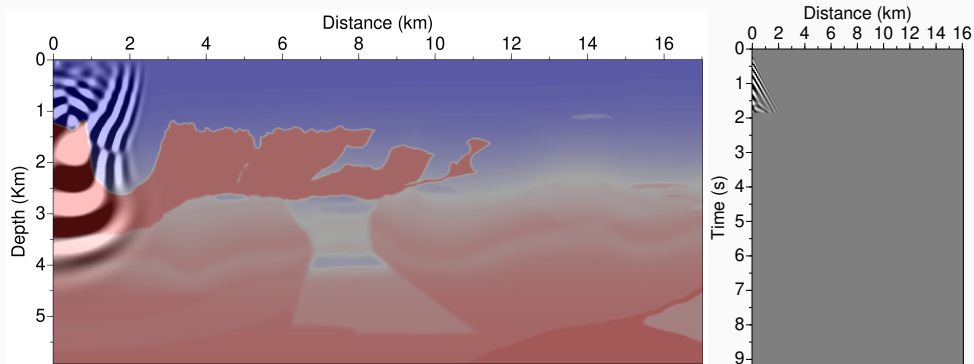
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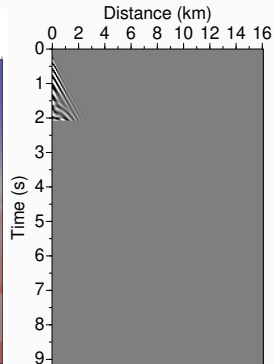
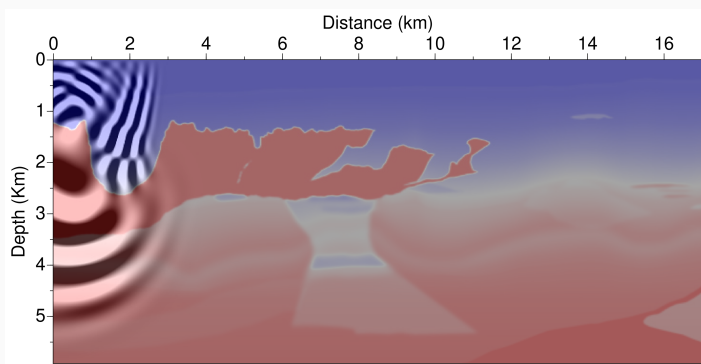
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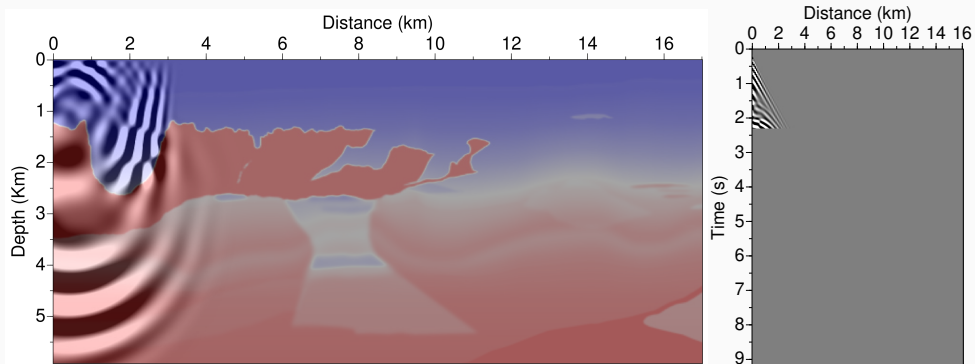
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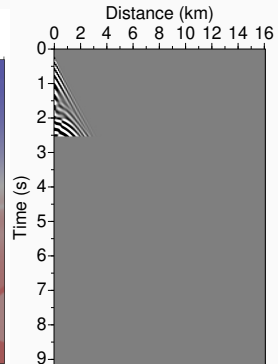
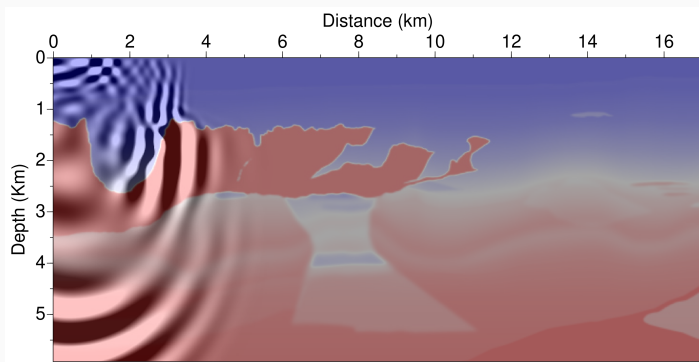
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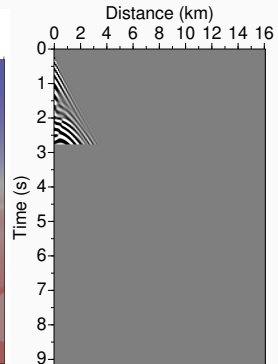
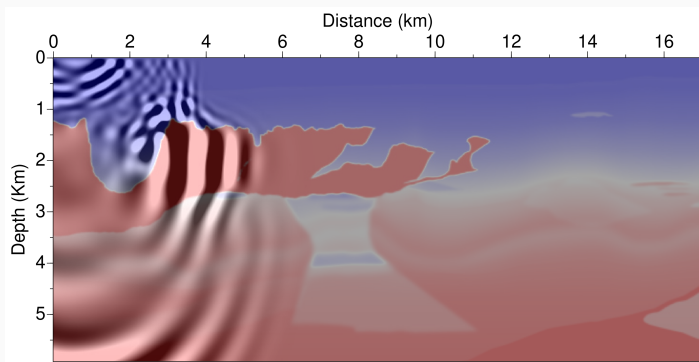


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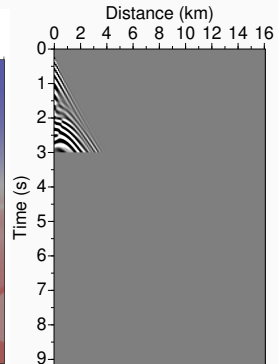
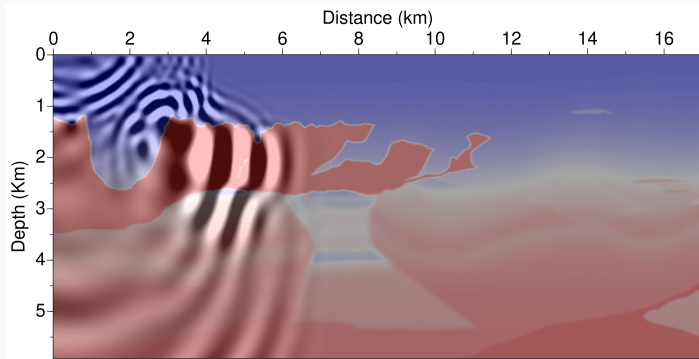




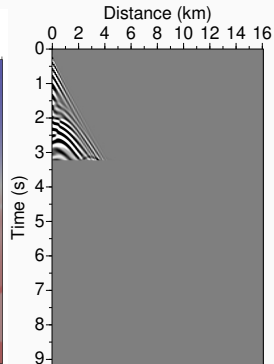
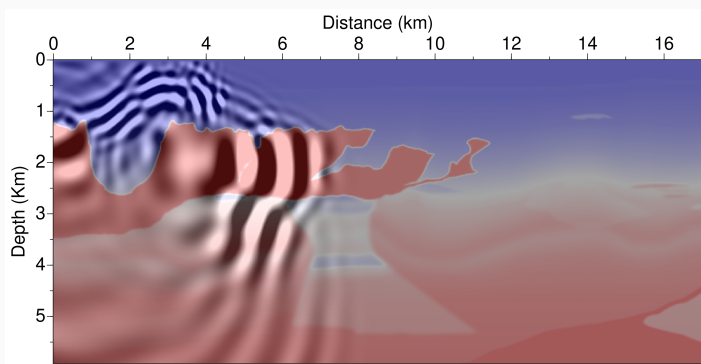
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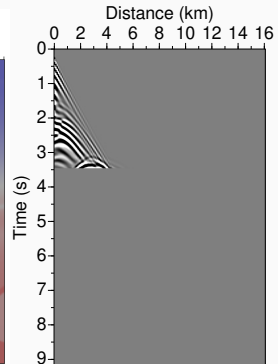
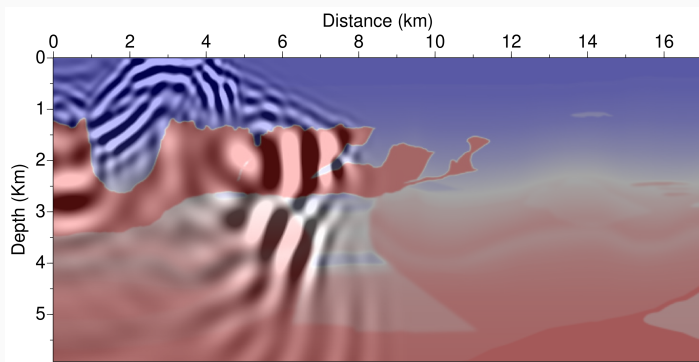
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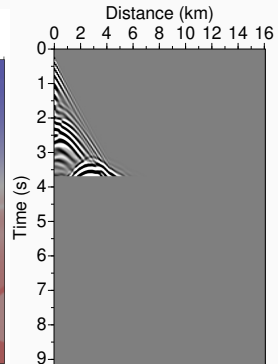
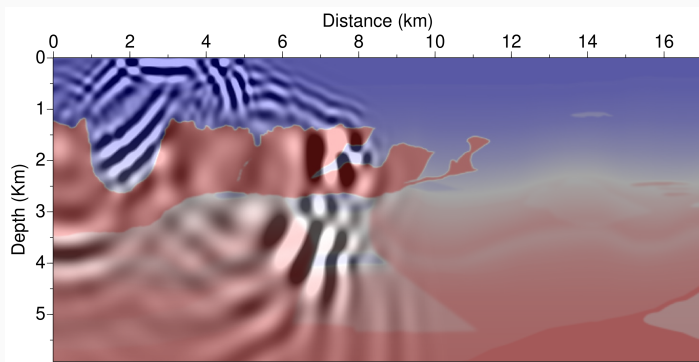
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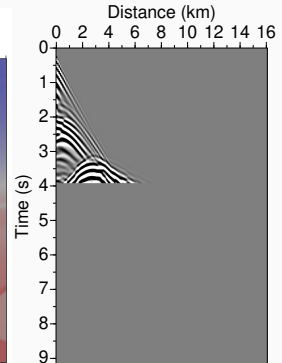
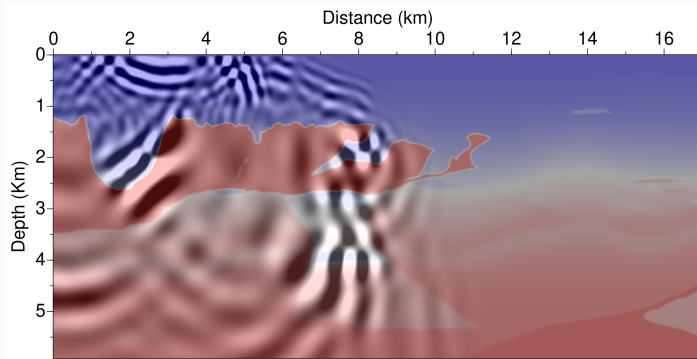
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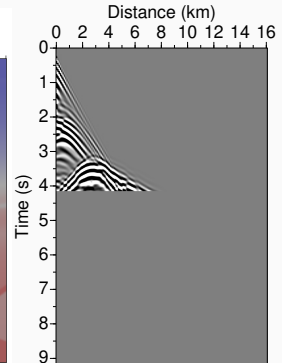
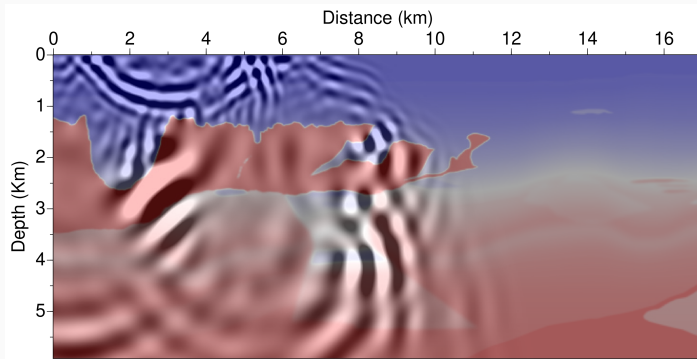
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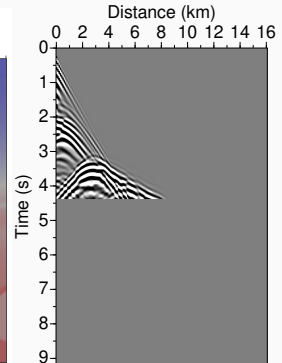
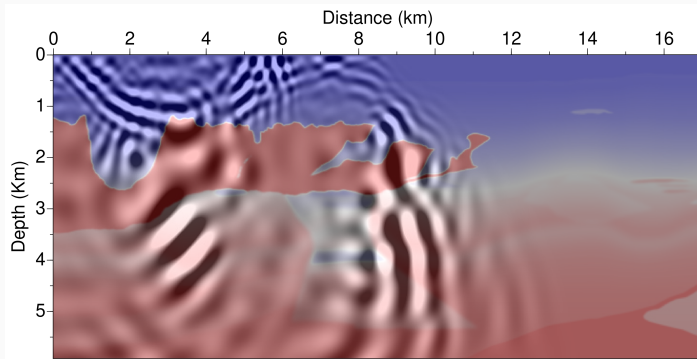
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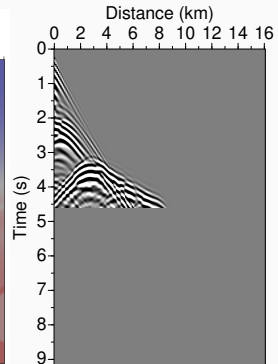
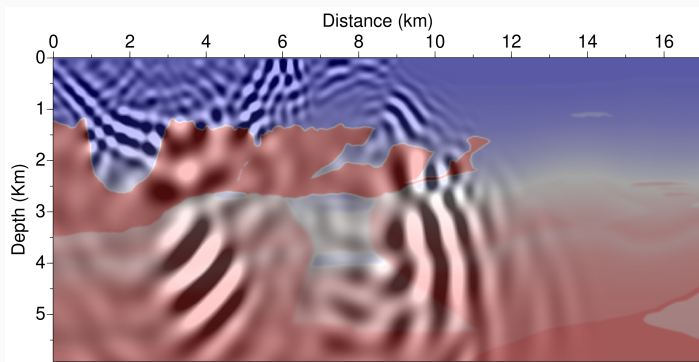


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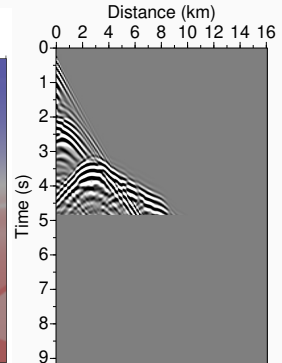
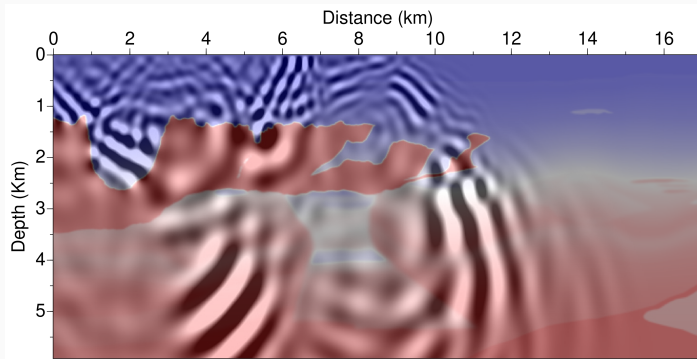




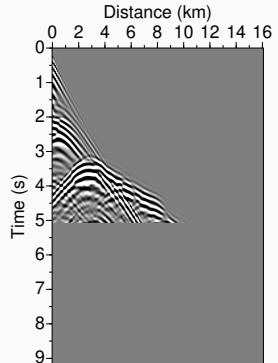
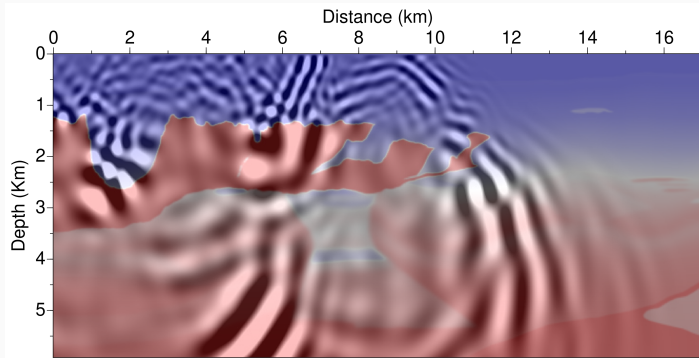
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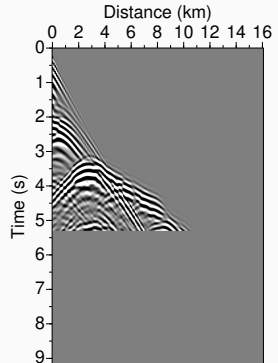
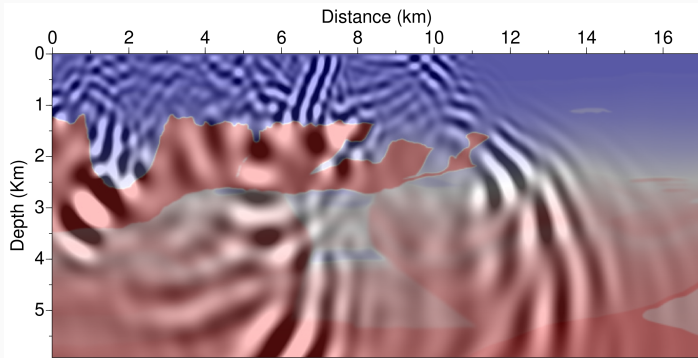
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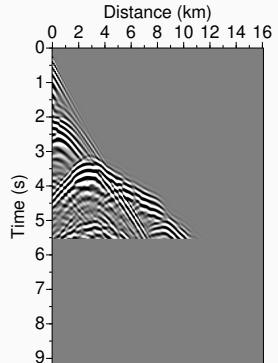
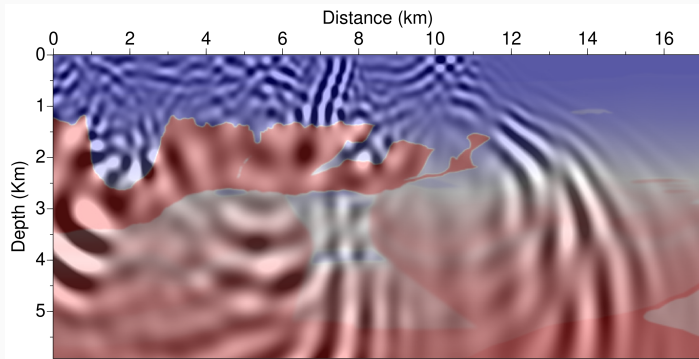
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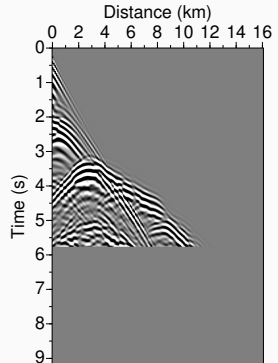
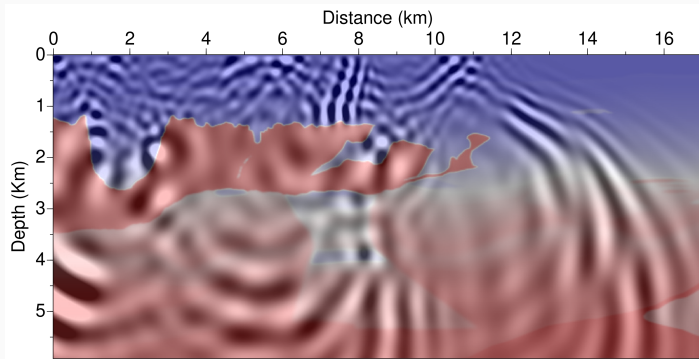
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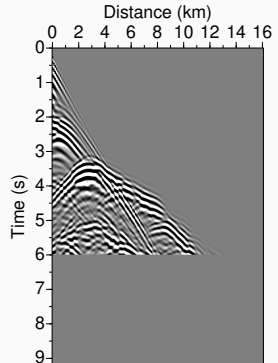
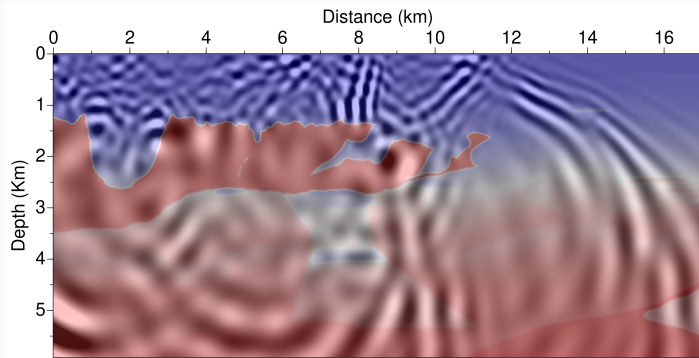
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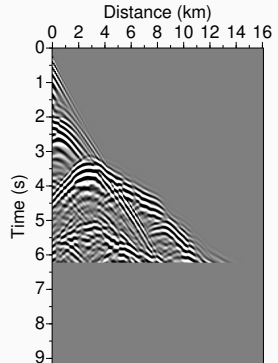
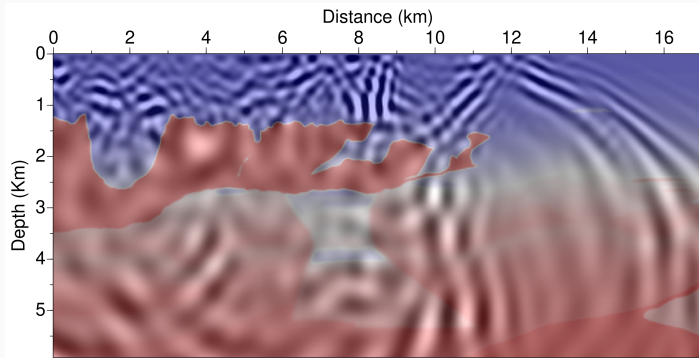
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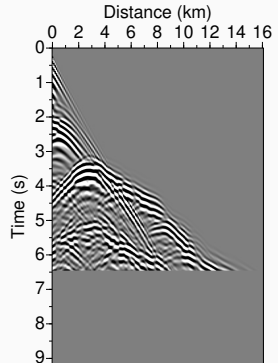
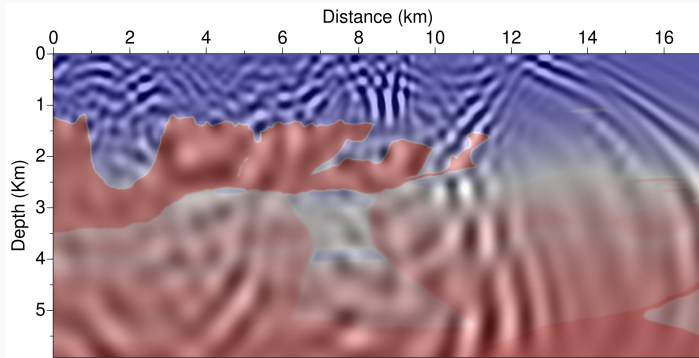


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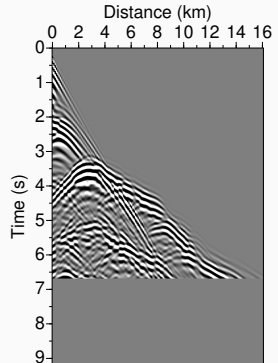
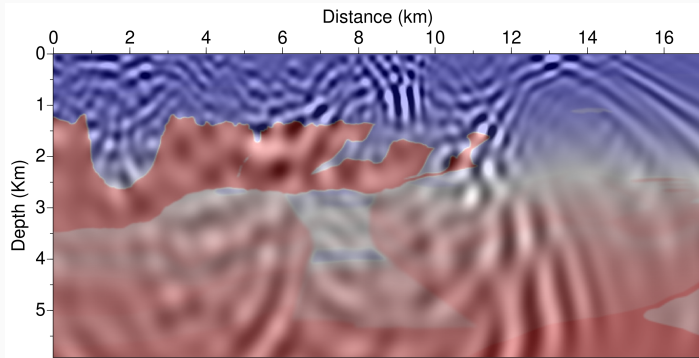




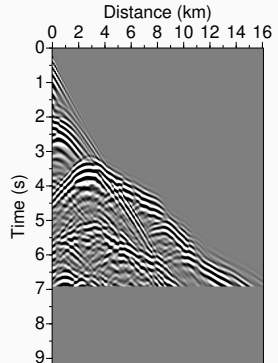
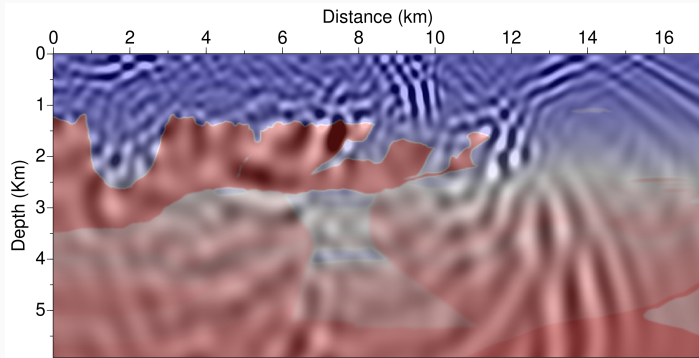
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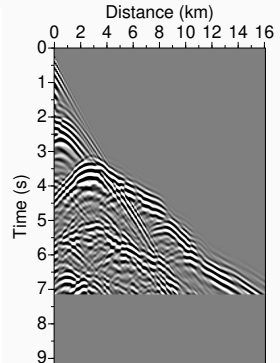
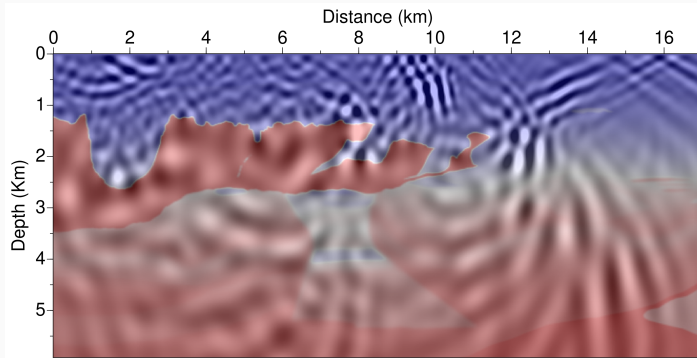
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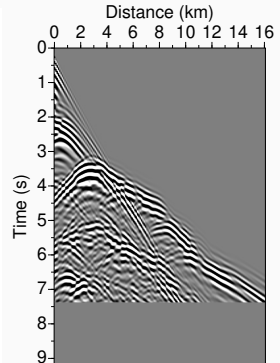
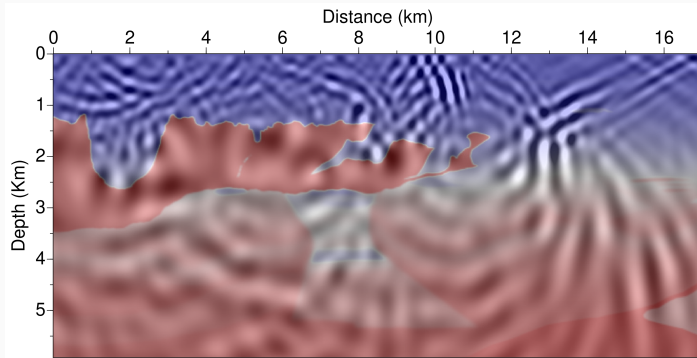
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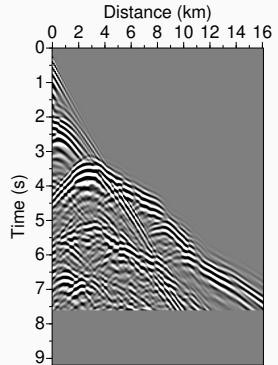
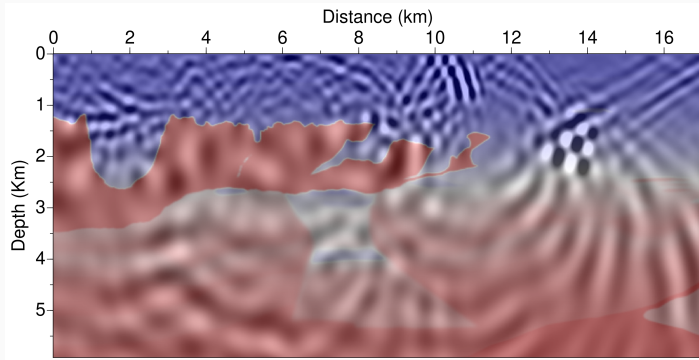
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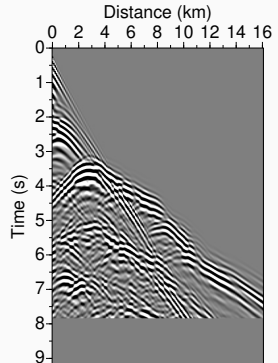
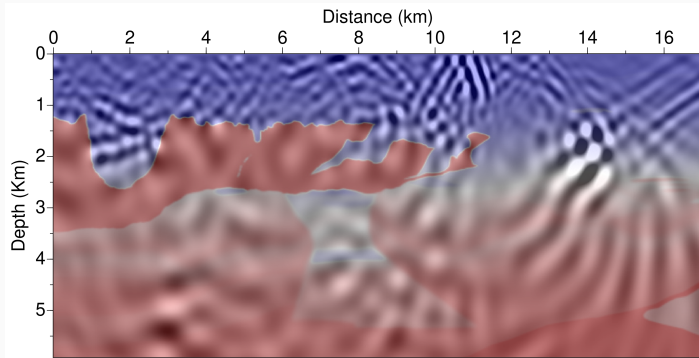
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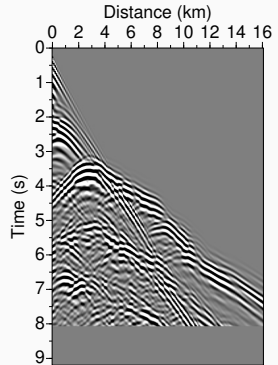
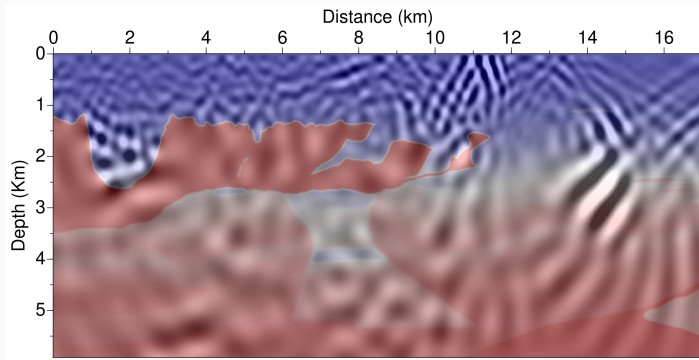
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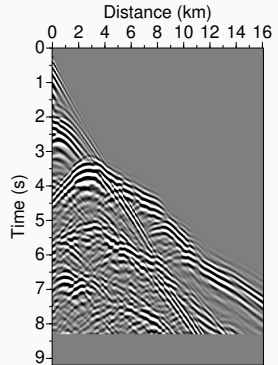
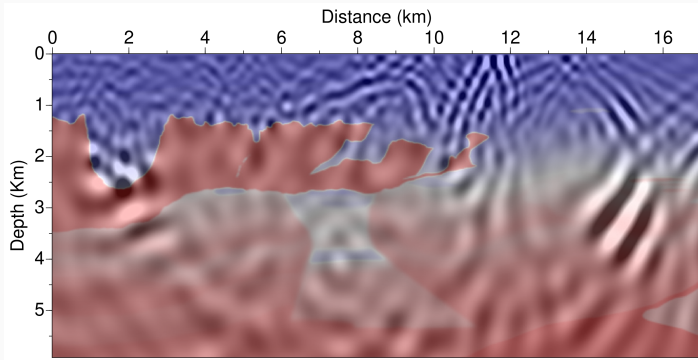


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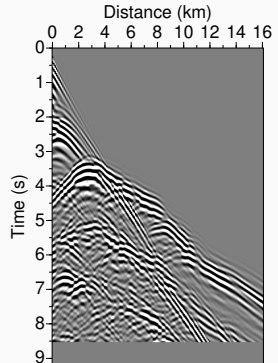
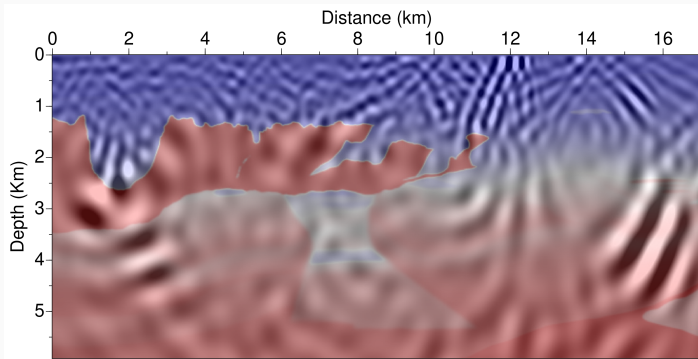




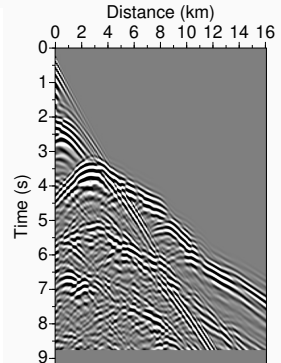
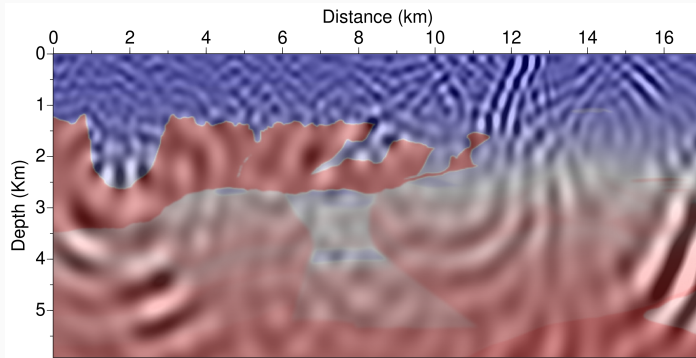
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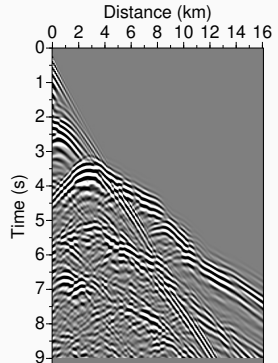
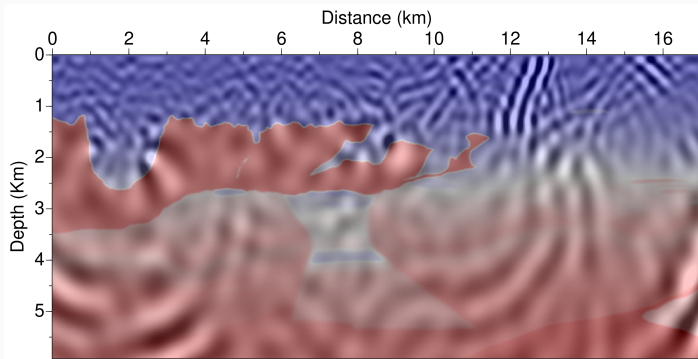
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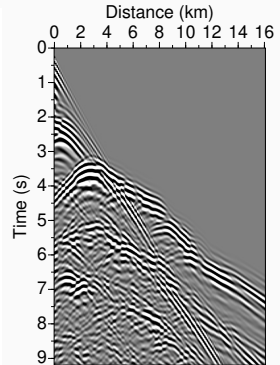
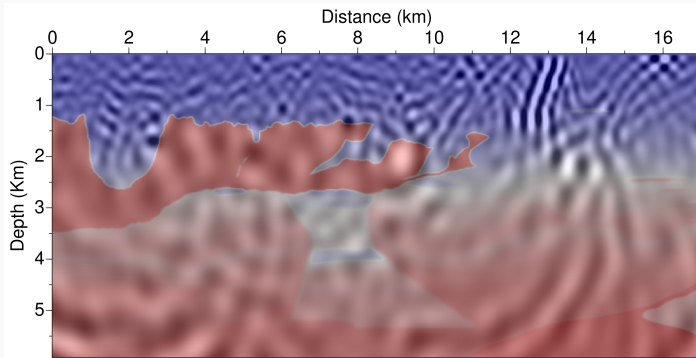
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## Introduction

Geophysical imaging: to do what?

Generalities on Inverse Problems

Seismic data

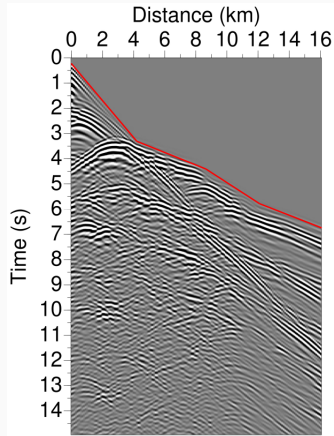
A first glance at seismic inversion methods

Full waveform modeling

Building the wave equation

Heterogeneity, anisotropy and attenuation

Full waveform inversion



Same seismogram as in Figure 22 with first-arrival travel time denoted by the red line.

Inverse tomography problem

$$d_{obs} = t_{obs}(x_s, x_r), \quad m = v_P \quad (5)$$

where  $t_{obs}(x_s, x_r)$  denotes the picked travel times from source  $s$  to receiver  $r$ , and  $v_P$  is the pressure wave velocity.



Least-square first-arrival travel time tomography

$$\min_{v_P} \frac{1}{2} \|t_{cal} - t_{obs}\|^2 + \eta R(v_P), \quad t_{cal} = g(v_P). \quad (6)$$

Idea: replace the forward modeling operator  $g(m)$  by a full wave modeling solver, and to compare the resulting synthetic data to the full observed data  $d_{obs}(x_s, x_r, t)$ . The FWI problem is thus formulated as

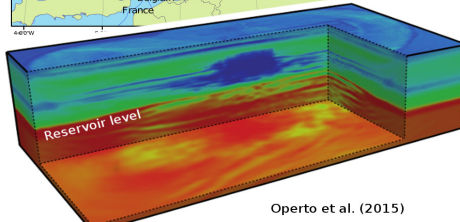
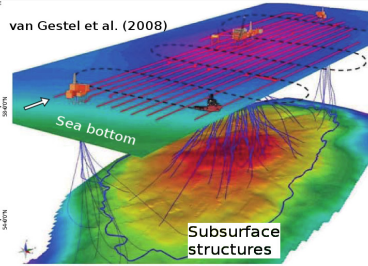
$$\min_m \frac{1}{2} \|d_{cal} - d_{obs}\|^2 + \eta R(m), \quad d_{cal} = g(m) \quad (7)$$

# 3D acoustic FWI of the Valhall data (Operto et al., 2015)

## Valhall target



van Gestel et al. (2008)

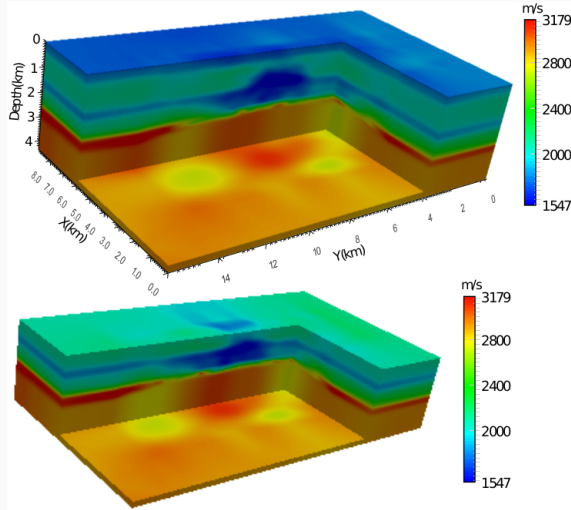


Operto et al. (2015)

- Area:  $\sim 145 \text{ km}^2$
- $\sim 50,000$  shots at 5 m depth
- $\sim 2,300$  receivers (4C) at 75 m depth (12 OBCs)
- Reservoir level  $\sim 2.5 \text{ km}$  depth
- Frequency-domain strategy : time-harmonic

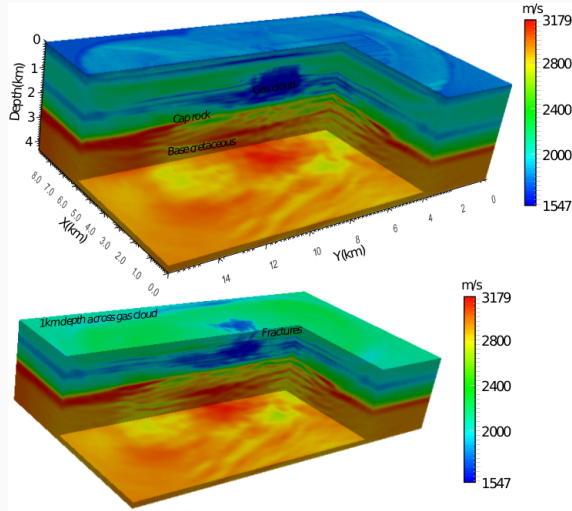
# 3D acoustic FWI of the Valhall data (Operto et al., 2015)

## Initial model



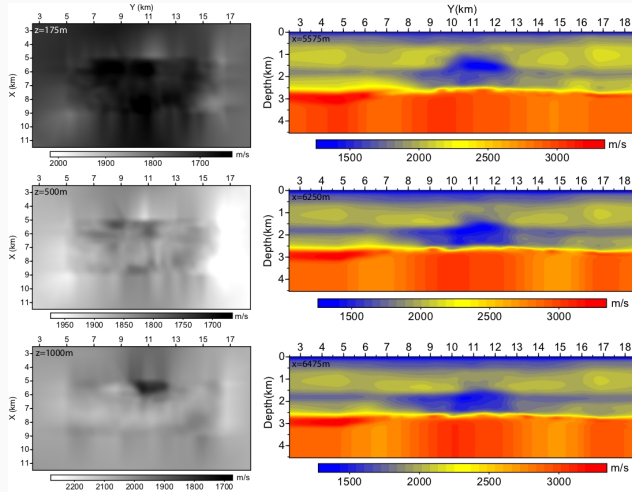
# 3D acoustic FWI of the Valhall data (Operto et al., 2015)

## Final FWI model



# 3D acoustic FWI of the Valhall data (Operto et al., 2015)

## Horizontal and vertical 2D slices: initial model



# 3D acoustic FWI of the Valhall data (Operto et al., 2015)

## Horizontal and vertical 2D slices: final FWI model

